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Chapter 5

CLIMATE CHANGE

APPROVED ASSESSMENT REPORT for the Grey Sauble Source Protection Area

October 15, 2015

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**APPROVED ASSESSMENT REPORT
for the
Grey Sauble Source Protection Area**

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5.0 CLIMATE CHANGE

5.1 What is Climate Change?

Climate can be defined as the meteorological conditions that are characteristic to a particular region, including temperature, precipitation and wind. The classical period for averaging these variables is 30 years, according to the World Meteorological Organization.

In a wider sense, the climate system refers to a highly complex system consisting of five major components: the atmosphere, the hydrosphere, the cryosphere, the land surface, and the biosphere, as well as the interactions between them. The climate system evolves over time through its own internal dynamics and under the influence of external forcing such as volcanic eruptions and variation in solar radiation (Intergovernmental Panel on Climate Change, AR4 WG1 Annex I).

Climate change refers to the change of the mean and/or the variability of statistical properties that characterize the climate system. To be considered an actual change in climate, and not simply a natural variation, the change must be persistent and measurable over time (Environment Canada, 2008). Three drivers for climate change have been distinguished. The first is natural internal processes, such as the evolution of land cover, major volcanic eruptions that alter the earth's atmosphere and planetary oscillations that include El Niño and the North Atlantic oscillation. The second relates to natural forcing that is external to the earth's climate system, which is mostly the intensity of solar radiation. The third, and most recent, driver is anthropogenic influences that change the composition of the atmosphere (especially by emitting greenhouse gases), the river systems, global nutrient cycles, and land use (Intergovernmental Panel on Climate Change, AR4 WG1 Annex I). The most relevant greenhouse gases are water vapour, carbon dioxide, methane, nitrous oxide, and ozone. These gases trap and reflect heat energy back to the Earth's surface (Lemmen and Warren, 2004).

As anthropogenic drivers of climate change, greenhouse gases are most significantly produced through the burning of fossil fuels for energy production, transportation and heating, as well as in agricultural activities and forest clearing (Lemmen and Warren, 2004). Lately, the emissions from peat oxidation after wetland drainage (Obenchain, 2004; Global Environment Centre, 2007) and livestock production (Food and Agriculture Organization, 2006) were recognized. The detection of climate change is now based on measurement values, especially in the polar regions.

Regarding causes of global warming, the formally agreed-upon statement of the Intergovernmental Panel on Climate Change (IPCC) concerning key findings and uncertainties states that:

“Most of the observed increase in global average temperatures since the mid-20th century is very likely due to the observed increase in anthropogenic GHG [greenhouse gas] concentrations. It is likely that there has been significant anthropogenic warming over the past 50 years averaged over each continent (except Antarctica).”

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(Intergovernmental Panel on Climate Change,
AR4 Synthesis Report: Summary for Policy Makers, 2007)

This statement was drafted by leading scientists and adopted by representatives of governments and intergovernmental organizations, including a delegation of the Canadian government.

5.2 Existing Data and Data Gaps

The impacts of climate change on the local hydrological cycle and its effect on water quality and quantity are poorly understood. While there is an understanding of climate change on the larger scale, very little is actually known about local impacts, particularly on a watershed basis. Any model-based research is limited by data availability, so better observational climate and hydrological data is needed to complete inventories and local scale data sets (Bates et. al., 2008).

The IPCC released their Fifth Assessment Report in 2013 and 2014. It provides an update on the current state of climate change and the science behind it. The complete report consists of three Working Group reports and a Synthesis Report. The Synthesis Report will be finalized on October 31, 2014. The Working Group I report is entitled “The Physical Science Basis” and was released in 2013. It contains a detailed assessment of climate change including chapters on sea level change, biogeochemical cycles, clouds and aerosols, and regional climate phenomena. Extensive information was taken from models and includes climate projections. The document also includes a new comprehensive atlas of global and regional climate projections for 35 regions around the world. The Working Group II report is entitled “Impacts, Adaptation and Vulnerability” and was released in 2014. It takes into consideration both human and natural system vulnerability, exposure and adaptation as well as impacts and future risks of climate change. The Working Group III report is entitled “Mitigation of Climate Change” and was released in 2014. This third report assesses the risks and ethics of climate change mitigation policies on all levels and for all sectors, including technological, economic and institutional (IPCC, 2013).

5.2.1 Global Scenarios for Climate Change and Model Intercomparison

As part of the regional chapter of the latest IPCC report, climate change was assessed for the North American continent and its five sub-regions using pre-defined scenarios for the development of humankind over the next century (Intergovernmental Panel on Climate Change, 2001). The leading global climate models were run under a rapid but balanced growth emission scenario (Intergovernmental Panel on Climate Change, 2001: A1B) and evaluated for large sub-regions. For the northern portion of Ontario, the corresponding sub-region is East Canada, Greenland and Iceland. For the southern portion of Ontario, the sub-region is Eastern North America, which includes this Source Protection Area (SPA) (Intergovernmental Panel on Climate Change, WG1-Ch. 11.5). These sub-regions are evaluated, and climate change between the current climate and the climate at the end of this century are computed, and results are compared between all climate models, which is referred to as multi-model intercomparison.

All models consistently show an increase in precipitation in the northern sub-region, with a model median value of 26% for the winter months and a lower increase of 11% for the summer months. For the full sub-region, temperature is expected to rise by 5.9°C in the winter and by

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2.8°C in the summer, which is reflective of the global trend of the strongest warming in the higher latitudes. For the southern sub-region, models show an increase in precipitation of 11% for the winter months and nearly no change (+1%) for the summer months. For the full sub-region, temperature is expected to rise by 3.8°C in the winter and by 3.3°C in the summer. The north will experience extreme seasons that are warm and wet, and the south is projected to have more frequent warmer seasons (Intergovernmental Panel on Climate Change, AR4 WG1, Ch. 11.5.3). However, the spatial resolution of global circulation models can only give limited insight into the climate development of Southern Ontario. This is mainly due to the fact that the sub-region is not representative for this most-southern end because of its unique location between the Great Lakes, which has a strong climatic imprint locally.

5.2.2 Regional Models and Scenarios

Within the Canadian Regional Climate Modelling and Diagnostics Network, the Canadian Regional Climate Model (CRCM) was developed for the North American domain. This model has a more detailed representation of processes, such as sea ice, clouds and mesoscale atmospheric convection schemes, than the global climate models, as well as a higher resolution and description of unique geographic features such as the Great Lakes (Goyette, 2000; Music and Caya, 2007; Plummer et al., 2006). Model quality for lake ice cover still remains difficult (Martynov, 2009).

The Canadian Centre for Climate Modelling and Analysis provides online access to climate data from this model with a 45 km horizontal grid-size mesh. The model uses observed data for the years 1961-2000 and assumes the rapid and heterogeneous growth emission scenario for the years 2001-2100 (Intergovernmental Panel on Climate Change, 2001: A2; Canadian Centre for Climate Modelling and Analysis, 2008). Recently, a new generation of the Canadian regional change model, GEMCLIM 2008, was presented. GEMCLIM was developed within the Meteorological Service of Canada using its modelling framework Global Environmental Multiscale (GEM).

A second attempt to model the regional impact of climate change for North America was developed in the United States. An impact report for the U.S. was published in 2009 by the United States Global Climate Change Research Program (USGCRP). It combines model outputs from the National Center for Atmospheric Research, the National Oceanic and Atmospheric Administration, the Geophysical Fluid Dynamics Laboratory, and the NASA Goddard Institute for Space Studies (Climate Change Science Program, 2008). The data is available online (Program for Climate Model Diagnosis and Intercomparison, 2009) and also from regional models (North American Regional Climate Change Assessment Program, 2009). The USGCRP report contains chapters for the state of Michigan, which can also be inferred for Southern Ontario due to geographic similarities.

Model projections for Michigan were evaluated under various emission scenarios, both for mid-century (2040-2059) and end-of-century (2080-2099) climate. In Michigan, summer average temperatures and precipitation are expected to feel progressively more like the summers currently experienced in the south and west towards the middle of the North American continent. In the year 2100, assuming the highest emission scenario, Michigan will experience a climate that corresponds to Northern Texas today. The models project that heat waves will be more

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frequent, more severe and longer lasting. However, both the frequency of hot days and the length of the heat-wave season will be more than twice as great under a scenario of higher greenhouse gas emissions than under a scenario of lower emissions (United States Global Climate Change Research Program, 2009). Other areas and sectors affected are urban areas (intensified heat island effect), flood control and stormwater systems, human health, energy, ecosystems, and the agricultural sector.

5.2.3 Existing Local Observation Data (Climate Stations)

Environment Canada's Atmospheric Environment Service (AES) provides daily climate data from climate stations in the form of hourly precipitation and temperature records (see Chapter 2: Watershed Characterization of this Report).

As is typical for such time series data, gaps exist for some time intervals and methodological inconsistencies have contributed to imprecise data records. Using statistical methods, data gaps were filled, methodological inconsistencies addressed (Schroeter and Associates, 2007) and continuous daily time series data for the period from 1950 to 2006 were created for five data points in the Grey Sauble SPA.

However, when evaluating climate data over this temporal interval, the impact of long-term climate change is convoluted with two natural large-scale oscillation patterns that cause decade-scale variability: the Arctic Oscillation/North Atlantic Oscillation and the El Niño-Southern Oscillation (Zhang et. Al., 1997; Latif and Barnett, 1994; Intergovernmental Panel on Climate Change, 2007 WG1, Ch. 9). Further, the eruption of the Agung volcano in 1963 started a decade of global temperature depression (Intergovernmental Panel on Climate Change, 2007 WG1, Ch. 9.4.1.2). If pulled together, the 1950s are typically characterized by elevated temperatures, while the 1960s are significantly cooler as evidenced by precipitation patterns that are associated with cooler temperatures. The stated reasons limit the analysis of data from local climate stations.

5.2.4 Trends in Local Climate Data

A project was completed by Huron Geosciences in 2011 to assemble, graphically display and analyze available meteorological data within the Source Protection Region (SPR) in order to give guidance to the SPR on providing services to their communities and member municipalities.

Overall, the trends indicate that total annual precipitation has increased over the 1950-2006 period. Temperatures have been increasing throughout the SPR with the exception of the Chatsworth station, which has a minor decrease. The increase in temperatures is largely the result of an increase in daily minimum temperatures.

Heat units are up across the SPR, as are the number of days where maximum temperatures have exceeded 30°C, which is indicative of warmer summer seasons. A marked decrease in the days where maximum daily temperatures do not exceed 0°C for the watershed is noted throughout the SPR, and is likely the result of a reduced frozen period and increased frequency and duration of winter melting events.

The study concluded that the impacts of the observed trends on drinking water sources will be low. In this SPA, the municipal systems are reliant on either deep bedrock aquifers or Lake Huron for their water supplies. From a water quantity perspective, increases in precipitation

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would lead to increases in the supply to these sources. The study highlights the increases in precipitation documented in the study when compared to the overall storage of these two systems.

Temperature increases could have a long term impact on water levels within Lake Huron, as they point to increases in evaporation (and notably, increased local precipitation). In particular, increased lake temperatures could lead to less ice cover on Lake Huron during the winter months and more winter lake-effect precipitation.

Increasing amounts of precipitation coupled with increases in intensity of precipitation are often associated with increases in soil erosion and ultimately on water quality of riverine surface water systems. In this SPA, there are no municipal systems that are exploiting riverine surface systems, and impacts on the overall quality of water in Lake Huron are buffered by the large volume of water in the system.

5.3 Expected Impacts of Climate Change on the Water Cycle in Southern Ontario

5.3.1 Introduction: Climate Change and Water

Climate change shows profound impacts on global and regional water systems, which are integral parts of the climate system. The increase of heat energy directly impacts the amount of evapotranspiration, atmospheric convection and precipitation that is received in any region, with feedback on land use (Intergovernmental Panel on Climate Change, 2008).

Precipitation is the condensation of atmospheric water vapour as rain, snow or other forms, if air is over-saturated with water. One cause is the addition of moisture to the air through evaporation from soils, from water surfaces or through precipitation/virga from higher atmospheric levels. Another cause is the increase of relative humidity by cooling and/or adiabatic expansion, for example, in upward movement of air during convection. Warmer climate impacts temperature and evapotranspiration, which affects the amount of water that air can transport and on temperature gradients that cause air convection. As a result, the amount and intensity of precipitation are affected, as well as its variability in time and space.

Evapotranspiration refers to the amount of water that is transported into the atmosphere from soils, water surfaces and plants. It depends on the amount of water that the air can transport, on temperature and relative humidity, as well as on the amount of moisture on the ground. In many regions, additional precipitation and warming will intensify evapotranspiration (Bates et. al., 2008).

If precipitation is more intense, larger portions of this water cannot be intercepted by plants or infiltrate into soils. Thus, intense precipitation creates more runoff into rivers and causes erosion of soils (de Loë and Berg, 2006), which can potentially reduce the amount of water recharging local aquifers.

Over the year, warmer temperatures will shift the form of precipitation from snow to rain and decrease the storage of snow in winter. Warmer summers will be drier and winters will be wetter.

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The spring thaw will probably occur more rapidly and cause more intense river runoff. Together, these changes are expected to cause more extreme flooding, landslides, coastal erosion, and droughts (de Loë and Berg, 2006).

The water cycle is not only modified because of climate change, but also by other human factors that influence the hydrologic cycle quantitatively as well as qualitatively. Urbanization and urban sprawl increase the sealing of soils and change the surface albedo. River modifications and wetland drainage, chemical pollution, and changes in the nutrient balance are other relevant factors (Aerts and Droogers, 2004).

In Southern Ontario, many relevant hydrological processes are closely linked to annual accumulations of snow in winter and melt in early spring. These processes include river runoff, the accumulation of moisture in near-surface soils and aquifer recharge (Rush, 2004).

The west side of the Bruce Peninsula is under the influence of Lake Huron and lake-effect snow, especially in winters with warmer Lake Huron water temperature and thus reduced ice cover. Lake-effect snowfall occurs if cold west/northwest winds travel over a significant stretch of open, warm water. This causes areas to the lee of Lake Huron to experience increased winter snowfall (Saugeen, Grey Sauble, Northern Bruce Peninsula Source Protection Region, 2008a). This phenomenon is linked to the migration tracks of mid-latitude cyclonic regimes, such as Alberta Clippers and other storms from the northwest. While global warming already seems to have increased the number of years with reduced ice cover on Lake Huron, little is known about how the decreased temperature gradient caused by arctic warming will impact these cyclonic patterns.

5.3.2 Great Lake Impacts

Water levels within the Great Lakes are regulated to a certain degree at the outflows of Lake Superior and Lake Ontario (Schertzer et al., 2008; Tupman, 2004). Several diversions exist throughout the basin. However, climate is the dominant factor affecting lake levels (Changnon, 2004). Lake Huron is not regulated; therefore, the impact of diversions is considered negligible in comparison with climate factors in controlling lake levels (Changnon, 2004).

The impact of climate change on the water budget of the Great Lakes depends on the relationship between precipitation and evaporation at the lake surface, on flow of surrounding rivers into the lakes and on other factors, such as groundwater inflows. In the past 150 years, annual average water levels in the Great Lakes have fluctuated up to 180 cm. Lake level data can be found on the Environment Canada website (<http://www.on.ec.gc.ca/water/levels>). In 2001, Lake Superior was at its lowest level since 1925 and Lakes Michigan and Huron were at their lowest levels since 1965. Low water levels reflect substantial loss of water volume in the Great Lakes system. Since this time, water levels in Lakes Huron and Michigan have rebounded significantly.

The exact balance of these processes is not understood well, in part because it is difficult to measure precipitation and evaporation over the Great Lakes and because no continuous measurement stations exist. Analysis of long-term regional climate data suggests that precipitation accounts for 55% of the variability in lake levels, with temperature accounting for 30% (Changnon, 2004). For the 1997 to 2000 period, temperature increases may be the primary

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cause of low water levels (Assel et al., 2004). In the long-term, net basin supplies will decrease with climate change. Lake levels will also decline due to increased evaporation and timing of precipitation (de Loë and Berg, 2006), which will be most pronounced in Lakes Michigan and Huron, at 0.73 to 1.18 m by the 2050s (Mortsch et al., 2006). Work in the Maitland Valley SPA has shown increases in both totals and intensity of precipitation over the period from 1950 to 2006, and it is not known what impacts this will have on lake levels or whether those trends are applicable to the Grey Sauble SPA (Luinstra, 2009).

Great Lakes ice cover is the second major climatic factor. Since the 1970s, the ice cover on America's Great Lakes has dropped by 30% (Great Lakes Environmental Research Laboratory, 2009). Moreover, Traverse Bay in Lake Michigan did not freeze over from 1995 to 2000, the first period of five consecutive winters without freeze over in at least 150 years (Assel, 2003). The ice cover of coastal regions is not noted and deep regions of the lakes remain open more frequently (Assel, 2007). Evaporation causes less ice coverage in winter, when cold, dry air passes over the open lake surface, which increases the amount of lake-effect snow in Southern Ontario. Other potential impacts include increased wave erosion, more hazards for transportation and navigation due to low water levels, potential reduction in hydropower generation due to drought, and the need for increased shoreline protection (Canadian Climate Impacts and Adaptation Research Network, 2006). Low water levels may also have negative impacts on recreational uses and the local tourism industry.

The combined impacts of climate change on the Great Lakes basin are complex as the area is influenced by hurricane tracks from the Caribbean, arctic air masses and pacific air masses. The manner in which global climate teleconnections interact with each other – especially the El Nino Southern Oscillation, the Arctic Oscillation and the Pacific Decadal Oscillation (Rodionov et al., 2003; Bai and Wang, 2009) – dominate the severity of winters, ice cover and summer heat in the Great Lakes area. Ultimately, the impacts of climate change on these global patterns will determine our future climate.

5.3.3 Groundwater

The number of freeze and thaw events has a significant impact on runoff distribution in the spring. Even if all data analysis uncertainties are acknowledged (see section 5.4.1), climate observations for the period from 1950 to 2006 document changes in the climate over the past two decades. Average annual temperatures have increased modestly, yet the annual number of days with highs that exceed 30°C has increased. A more pronounced trend suggests that there are fewer days where temperatures remain below 0°C. This trend likely reflects a shorter frozen period or an increase in the number of freeze and thaw cycles (Luinstra, 2009).

Data and modelling indicate that winter months are most significant recharge times, peaking in late autumn and spring thaw (Saugeen, Grey Sauble, Northern Bruce Peninsula Source Protection Region 2008a; Prodonović, 2008). Further research is needed to validate these results.

Reduced snow cover and a shorter frozen period, together with faster thawing, are likely to diminish the quantity of recharge to local aquifers, which could have negative implications for private well owners that rely on shallow, unconfined aquifers for drinking water sources. This could lead to a shift from the exploitation of shallow aquifers to deep aquifers, which would

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result in additional demand for the latter (Lemmen and Warren, 2004) and have uncertain impacts in this Source Protection Area. Implications on fish populations from reduced groundwater discharge to rivers, which is relevant during the low-flow season in summer and autumn, is still unknown (de Loë and Berg, 2006).

5.3.4 Surface Water Flows

Surface water flows comprise the water in rivers, their tributaries and creeks and also includes water in reservoirs, wetlands and lakes. Surface water is derived from yearly snowmelt, base flow from groundwater, runoff from precipitation events, and other human intervention, such as dams and flood control measures.

IPCC multi-model intercomparison (see section 5.2.1) suggests that the annual amount of precipitation decrease around the Great Lakes is comparatively low, with a projected zero to five percent decrease. However, precipitation is projected to increase in winter and spring, and models and recent observations indicate that the intensity of precipitation events increases during the rest of the year (United States Global Climate Change Research Program, 2009; Luinstra, 2009) with a longer duration of low-flow periods (de Loë and Berg, 2006). The intense precipitation events seem to be more localized, causing local flash floods and erosion events. While such events are difficult to measure with the existing meteorological network because of their small spatial extent, their impact on local infrastructure can be profound (Pearson, 2009).

Earlier and lower spring freshets – the flow resulting from melting snow and ice – are expected for Southern Ontario (Lemmen, 2008). The increase in thaw events and flash floods has the following implications:

- maintenance of municipal infrastructure, such as roads and bridges
- sewage and water treatment facilities
- increased capacity demand on storm sewers and stormwater management systems
- dams, shore erosion control and flood control measures

The changing environmental conditions should be taken into account, especially for land use planning and for those investments that have a turnover time of several decades, which include sewer and drainage systems, water treatment plants, and flood and erosion control measures (Burton, 2008).

Within the Maitland Valley Conservation Authority, precipitation increases have been observed during the period from 1950 to 2006, which can be attributed to decreased ice cover on Lake Huron and the northerly migration of winter storm tracks for mid-latitude cyclones. Higher intensity rainfall events lead to increased erosion of gullies near the lakeshore, increased soil erosion and a deterioration of surface water quality (Luinstra, 2009).

5.3.5 Wetlands

Climate change impacts on wetlands are still insufficiently understood. Coastal wetlands are an important habitat for waterfowl as well as breeding and nursery areas for many fish. The lowering of Great Lakes water levels could modify or eliminate coastal wetlands and reduce their ecosystem functions (shoreline protection, erosion reduction, contaminant filtering) as reservoirs

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that absorb excess storm water (Lemmen, 2008) and as carbon storage (Obenchain, 2004). The impact of climate change must be evaluated in conjunction with other stressors, such as urban development and agriculture (Ministry of Natural Resources, 2009). Key wetlands are at risk, particularly those that are impeded from adapting to the new water level conditions by man-made structures or geomorphic conditions (Mortsch, 1998).

Inland wetlands, especially within this SPA, are closely linked to the groundwater and surface water cycle and are even more difficult to assess since they partly replenish groundwater aquifers and are partly supplied by groundwater. The understanding of climate change impacts on wetlands requires improved data on the role of Karst aquifers, surface water bodies, aquifers, and groundwater recharge. Indirect climate impacts due to changing human land use may also impact wetlands.

5.3.6 Other Implications

Impacts on maintenance costs for electricity and communications lines are expected. For example, in the Toronto area the cost of insurance as a result of flooded basements and buildings and extreme weather events is reported to have multiplied by more than 13 times from 1960 to 1999 (CAP, 2007). Some municipalities in the neighbouring Maitland Valley SPA reported an increase in road maintenance costs, due to premature deterioration of roads, after an increase in the freeze and thaw cycle (Pierson, 2009, Mills et. al., 2007).

5.4 Implications for Drinking Water Source Protection

5.4.1 Direct Impacts on Drinking Water Sources

The amount and availability of municipal drinking water in the Grey Sauble SPA is not expected to be significantly impacted by changes in climate in the next few decades. The largest impact is expected for private overburden wells and Karstic systems; however, the impact on these systems is not known (Saugeen, Grey Sauble, Northern Bruce Peninsula Source Protection Region, 2008b).

Most municipal drinking water systems for small urban centres rely on wells that retrieve water from confined bedrock aquifers, which are not expected to experience any significant reductions in water availability in the near future. Impacts are most likely in systems that rely on overburden wells, inland lakes or springs that are impacted by precipitation and flood events. Refer to chapter 7 for more detailed evaluation of potential climate impacts on municipal drinking water systems. Municipal drinking water systems that are exploiting Karstic aquifers may experience more frequent periods of contamination due to increasing intensities of rainfall and more frequent snowmelt events.

Great Lake intakes will be affected by the change of lake level, as well as by the change of the physical, chemical and biological properties of lake water. Most intakes in the Grey Sauble SPA are located in deep water so that, even employing a conservative estimate, no deterioration in the ability of the intakes to provide water can be reasonably projected. In the near future, lake levels are not expected to drop more than one metre (Luinstra, 2009). However, it should be noted that no standard exists for intake depth, which could provide guidance on optimal intake depth.

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Climate change may cause or trigger water quality issues and health issues in drinking water systems. These are related to the decrease of water quantity (decrease of dilution effects), new and invasive pathogens under a new temperature regime, and extreme events such as water intrusions caused by flooding and stormwater runoff. The role of transport pathways, such as abandoned or improperly decommissioned wells, may play a significant role. Especially for Karst bedrocks, the potential impact of increased rainfall intensity is poorly understood.

Finally, the overall impact of climate change on this SPA is intrinsically linked to the impacts on the local hydrologic cycle. Lake Huron is the dominant local driver of climate in the area; therefore, any impacts on it will have a profound effect on local climate patterns.

5.4.2 Indirect Impacts on Drinking Water Sources

With climate change impacting the water cycle, humans and ecosystems will adapt to their new environment. This adaptation of sectors such as agriculture, tourism, industry, and others (Field et al., 2008) will modify the water cycle indirectly.

With increasing variability of precipitation, the agricultural sector might increasingly rely on supplemental irrigation practices (International Assessment of Agricultural Science and Technology for Development, 2008). To mitigate the risk of yield loss, farmers may use either surface water (e.g., from retention ponds) or groundwater for irrigation water. Especially for crops that require a high level of input and have associated investment costs, either because of manual labour, energy, chemicals, or any other form, the high energy cost associated with groundwater pumping can be economical. Examples of activities where irrigation security requirements are highest and groundwater-based irrigation is most likely include contract farming, perennial plantations and vegetable farms.

With the increasing strength of climate change, the relevance of food security concerns will also impact the general farming patterns (International Assessment of Agricultural Science and Technology for Development, 2008b). While a shift toward an increasing share of local production of high value products that rely on groundwater irrigation can be expected, a great level of uncertainty remains.

The tourism sector may change and its water demands may significantly increase, especially if living conditions in nearby urban centres are less favourable than today during the hot and dry summer months. The changing living conditions in urban centres might change the water demand.

The quality requirements for drinking water are higher than for any other sector – for human consumption as well as for livestock. Under a scenario of climate change, the pressure on existing drinking water resources might increase significantly because other sectors increase demands. As a result, the provision of drinking water from clean groundwater sources can increasingly compete with the increasing demands from other sectors and require improved management.

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5.5 Knowledge Gaps

Groundwater has traditionally received less attention than surface water, despite its importance as a drinking water source worldwide. While groundwater has been monitored in the past, it has not been monitored and examined for the eventuality of a changing climate and the associated impacts. Such analysis would require an improved understanding of aquifer formation, sources and natural discharges, for example, into the Great Lakes. However, depletion levels and recharge rates have neither been modeled nor measured well, and existing models are based on limited data (Bates et al., 2008). More study is needed on core processes such as recharge, aquifer discharge and in-aquifer water movement, before confidently assessing the impacts of climate change on water resources at a watershed scale.

With respect to surface water flow, local trends in climate data suggest that precipitation is becoming more intense and localized, which has significant implications for flood management. Due to their local nature, an early warning system for such intense events requires a denser measurement network than exists today.

The impact of climate change on mid-latitude cyclones on this region is profound. In winter, these cyclones carry precipitation from outside the Great Lakes basin into the area. In addition, lake-effect snow heavily impacts Southern Ontario, and is highly sensitive to the extent of ice cover on Lake Huron. While data indicates that this ice cover is changing dramatically, neither the impact of climate change on the storm track of cyclones, nor the changes in lake-effect precipitation, are well understood (Luinstra, 2009).

A significant, indirect impact of climate change is expected, if other human sectors change their water consumption behaviour, such as the agricultural sector or the tourism sector. This human adaptation to climate change has the potential to impact water quantity and quality. However, very little research is available to estimate the magnitude of these potential impacts.

Further, potential climate change impacts on Karstic aquifers within this SPA remains a core knowledge gap. There is a general lack of understanding of these Karstic systems and their interaction with the local climate.

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